

(ii) Depression storage:-

தீர்த்த வளைநிலக்கீழ்
(compact soil)

This is one of the catchment area generally many depressions are shallow depths and of varying size and shape

The precipitation (படி) will be falling on land surface first water runs towards the depression storage and filling it

The storage water in these depression storage partly evaporates and partly infiltrates into the ground

The following relationship may be used for computing the depression storage

$$V_d = k [1 - e^{-Pe/K}]$$

where

V_d = volume of water stored in surface depression

Pe = Rainfall excess (volume of rainfall in excess of infiltration and interception)

k = depression storage capacity of the basin

The following are the factors that affect depression storage

- (i) Land form (நில வடிவமுறை வகைகள்)
- (ii) Soil characteristics (நில சிறைகள்)
- (iii) Topography (நில மூலக்கூறுகள்)
- (iv) Antecedent precipitation index and (முறையில் விடுவீசு)
- (v) Man made disturbance etc

~~etc~~ terrace farming (நிலங்கள் பயிரிடுதல்)

(iii) watershed leakage:

watershed leakage may be defined as flow of water from one basin to another basin, (or) from one basin to the sea through major faults, fissures

(நிலங்களின் இடமாக விடுவது)

ஒரு பகுதி யில் நிலங்கள் காலி தான் என்ற உரிமையை கொண்டும் நிலங்களில் விடுவது என்று அழைக்க வேண்டும் என்று அமைக்க வேண்டும் என்று அமைக்க வேண்டும்

Major faults, fissures may be formed due to the catchment water will be moving downwards ~~near the water table~~. At time will be formed fissures (ஏரங்கு)

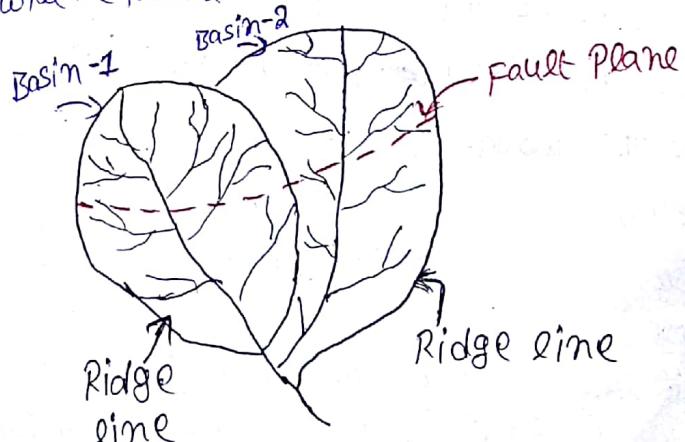
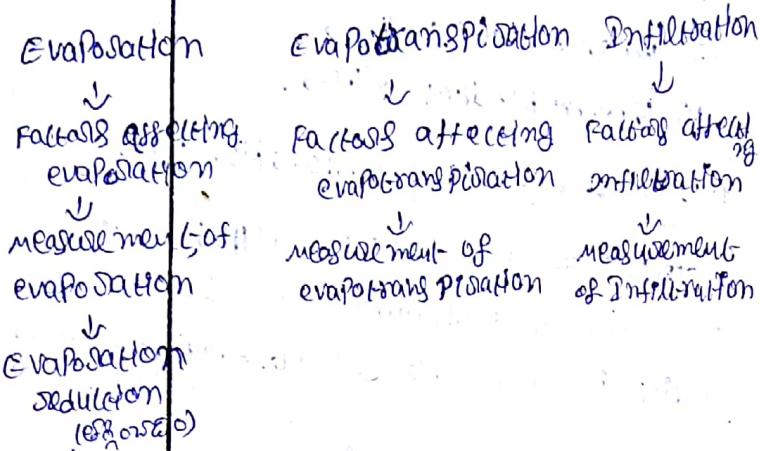


Fig 4.29 watershed Leakage

UNIT-II

In this chapter we can study only 3 topics



Evaporation:

- * Evaporation is the process in which liquid form of water changes to gaseous state at the force of heat energy.
- * It is a continuous process in natural porcesses by which a substance (solid, say) changes from liquid to gaseous state.
- * The main source of evaporation is the solar radiation.

* The loss of evaporation is high, may be 90% of annual precipitation will be evaporated.

Factors affecting evaporation:-

- (1) Nature of evaporation surface.
- (2) Area of water surface
- (3) Depth of water in the water body.
- (4) humidity.
- (5) wind velocity.
- (6) Temperature.
- (7) Atmospheric pressure and (8) quality of water.

(1) Nature of evaporation surface:-

- * different evaporating surfaces like soil, bushland, forest area, houses and lakes
- * Black cotton soil will absorb the heat very fast so black cotton soils helps to evaporate the soil water faster comparatively than red soils

(2) Area of water surface:-

- * The evaporation loss directly depends upon the area of water surface
- * The water surface area is more the evaporation of water will be more
- * The water surface area is less the evaporation of water will be less.

(3) Depth of water in water body:-

- * Deep water bodies the water will be evaporated slower. comparatively shallow water bodies.
- * In shallow water bodies the water will be evaporated faster.

in summer while in winter season

(4) Humidity:-

- * If the humidity in the atmosphere is more, evaporation will be less.
- * If the humidity in the atmosphere is less, evaporation will be high.
- * Evaporation is inversely proportional to humidity.

(5) Wind velocity:-

- * wind removes over laying (top layer) vapour from a evaporating body.
- * This heat vapour will reflect behind the water body to absorb the vapour
- * thereby increasing the rate of evaporation

(6) Temperature of air:-

The air temperature will be moderate due to the fact that air will meet cold objects. The evaporation rate will be more.

The air temperature will be less the rate of evaporation will be less. In summer season the air will be very hot, so evaporation will be high. In winter season the air will be cold, so evaporation will be less.

(7) Atmospheric Pressure:-

Hot air don't have any weight so it will be flowing upper surface.

* on cold air having some weight so will be flowing in down stream of hot air & meay above the ground surface.

* on heavy cyclone is coming that a time high speed wind will come this wind will be flowing fast & meet any obstacle and then turn upside near hot air.

After that this hot air become cool because bottom side of cool air will meet upper side of hot air so become cool. so evaporation will be less.

* sometimes the upper surface of hot air meets snow maintaining that a time evaporation will be more.

(8) Quality of water:-

The water having any dissolved salts the water will be less evaporating.

Ex:- Sea water is less evaporating
River water is more evaporating

Measurement of evaporation

The following methods while using the measurement of evaporation.

(1) Measurement using evaporation pans

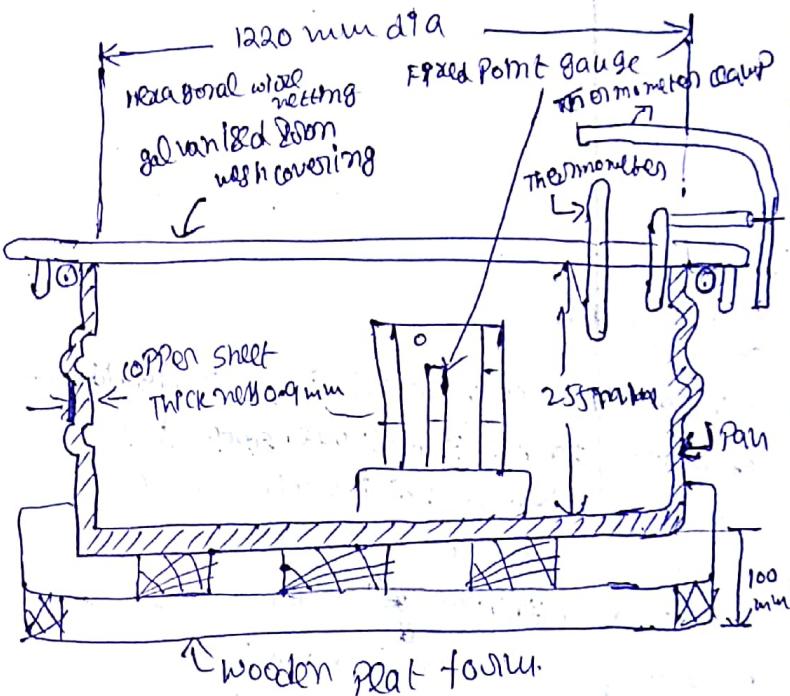
(2) use of empirical equations

(3) water budget method

(4) Energy budget method.

(1) Measurement using evaporation Pans

This is the most suitable method for the estimation of evaporation from large water bodies is that by measurement from evaporation pans.



1225 mm

* A Pan of 1220 mm dia and 255 mm deep and thick ness of 0.9 mm with copper sheet and this equipment covered with hexagonal wire netting of galvanised iron mesh covering.

* This water ~~because~~ ^{to use} of protecting birds.

* maintain the uniform temperature during day and night.

- * The pan is placed over a square wooden platform of 12.25 mm width and 100 mm depth (or) height so that circulation of air is possible all around the pan.
- * Water level in the pan is recorded by a point gauge.
- * Measurement is taken at least once a day by adding water to the pan is calculated experimental glass jar to bring the water level in previous position.
- * This gives directly how much of depth of water will be evaporated.
- * The evaporation of pan, evaporation of lake and evaporation of reservoir will be difference due to the following reasons.
 - (i) Depth of reservoir of pan above ground.
 - (ii) Colour of the pan
 - (iii) Height of the atm (8.1 m)
 - (iv) Heat storage and heat loss for capacity with respect to reservoir (0.7 m)
 - (v) Rain discharge.
 - (vi) Variation in vapour pressure, wind speed and water temperature.

Due to these reasons lake evaporation will be less than the pan evaporation.

The calculation of lake evaporation is as follows:

Lake evaporation = Pan coefficient \times Pan evaporation

+ The Pan coefficient for the Indian standard evaporation meter is around 0.8.

$$V_z = C \cdot z^{1/4}$$

The value of $V_{0.6}$ is computed by the following equation

$$\frac{V_{z_1}}{V_{z_2}} = \left[\frac{z_1}{z_2} \right]^{0.143}$$

where V_z = velocity at any height z

(a) Use of empirical equation:-

There are a large number of empirical equation for estimating the evaporation loss.

We consider the following two popular formulae

(1) May's formula (2) Rohwer's formula

(1) May's formula-

$$E = k_m (e_s - e_a) \left[1 + \frac{V_q}{16} \right]$$

where

E = evaporation

e_s = saturation vapour pressure at t_w .

t_w = water surface temperature

e_a = Actual vapour pressure at t_w

V_q = monthly mean wind velocity in m/h at height of $9m$ above the ground.

k_m = coefficient accounting for variability
other factors
 ≈ 0.36 for large deep water
 ≈ 0.50 for small shallow water

(2) Rohwer's formula

$$E = 0.771 (1.652 - 0.000732 \cdot P_a) (e_s - e_a)^{0.6}$$

$V_{0.6}$ = mean wind velocity in m/h at $0.6m$ height above ground level.

E , e_s and e_a have the same meaning as in May's formula.

P_a = mean atmospheric pressure (i.e.) barometric reading in mm of mercury

(3) Water budget (or) water balance method:

This method balances all the incoming, out-going and stored water on a lake(s) over a period of time, using the following reservoir equation.

$$\Sigma I - \Sigma O = \Delta S + E$$

Inflow + Outflow = Change in storage + Evaporation loss

$$\Sigma I - \Sigma O = \Delta S + E$$

$$E = \Sigma I - \Sigma O \pm \Delta S$$

The above equation can be generalized as under, taking all the factors of inflow and out-flow

$$E = (P + I_{sf} + I_{gf}) - (O_{sf} + O_{gf} + T) \pm \Delta S$$

where: P , precipitation.

I_{sf} = surface water inflow

I_{gf} = ground water inflow

O_{sf} = surface water out-flow

O_{gf} = ground water out-flow

T = transpiration loss, which may be neglected.

ΔS = change in storage

The above equation does not give accurate result because it is very difficult to measure I_{sf} and O_{gf} for a lake/reservoir.

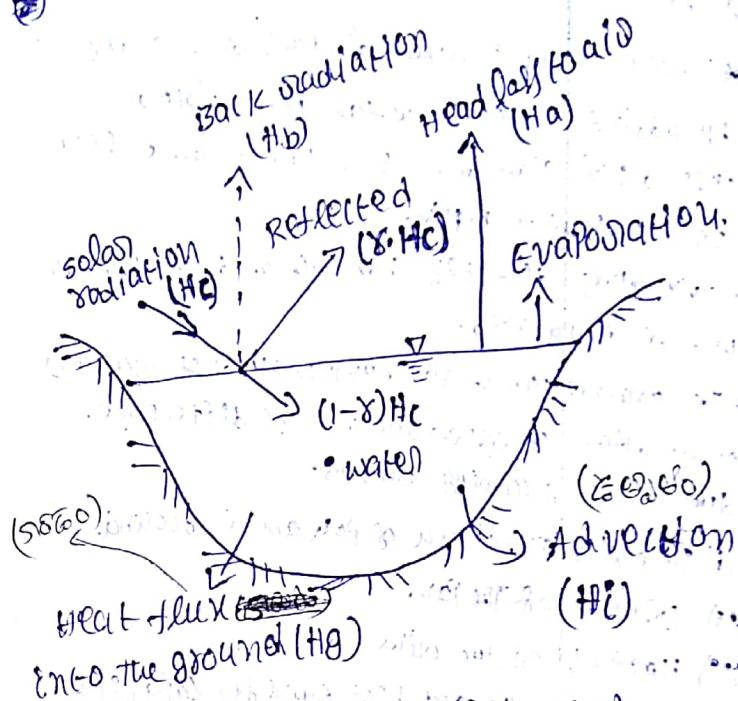
(4) Energy budget (or) energy balance method:

This method is based on the principle of conservation of heat, energy by evaporation

thus, the energy available for evaporation is determined by considering the incoming energy, out-going energy and energy stored on the water body over a known time interval.

Let us consider a water body, as shown in fig. The energy balance at the evaporating surface in a period of one day (say), is then given as

(i)



(ii) Net heat energy derived (Hn)

Net incoming energy

$$Hn = Hc(1-\gamma) - Hb$$

where

$Hc(1-\gamma)$ = The incoming solar radiation into a surface of reflection coefficient " γ " as shown in fig

The value of " γ " for water surface approximated 0.05; and for newly laid snow = 0.90.

Hb = back radiation (long wave) from water body

(iii) Out-going energy = $Ha + Hg + Hi$

where Ha = heat lost to air.

Hg = heat lost to ground

Hi = net heat conducted out of

the system by water flow (adverted energy)

(iii) Heat stored in the waterbody (Hs)

(iv) Heat used in evaporation

$$(He) = P_w \cdot La \cdot E$$

where

P_w = density of water

La = latent heat (quality of heat it is not visible)

of vaporisation of water.

E = evaporation in mm

using energy conservation law, we have

Total incoming energy = out-going energy
+ energy stored + energy used on evaporation

$$Hn = (Ha + Hg + Hi) + (Hs) + (He) \rightarrow ①$$

(v) All the energy terms in the above eqn. are in calories/mm/day.

Moreover, all the terms except Ha in the above eqn. can be measured or evaluated indirectly.

Ha can be estimated by using

Bowen's ratio (β), given by the eqn.

β = Heat stored formed water surface to the atmosphere (Ha)

Energy used on evaporation (He)

$$= 6.1 \times 10^{-4} \text{ Pa} \left[\frac{T_w - T_a}{e_s - e_a} \right] \rightarrow ②$$

where Pa = Atmospheric pressure in mm of Hg

T_w = water temperature in $^{\circ}\text{C}$

T_a = air temperature in $^{\circ}\text{C}$

e_s = saturation vapour pressure at T_w in mm of Hg.

e_a = air vapour pressure in mm of Hg.

Eqn ① and ② can further solved as $\rightarrow ③$

$$Hn - Hg - Hi - Hs = Ha + He \rightarrow ③$$

$$\frac{Ha}{He} = \beta \quad \text{we know that}$$

Adding 1 both sides

$$1 + \beta = 1 + \frac{Ha}{He}$$

$$1 + \beta = \frac{Ha + He}{He}$$

$$Ha + He = (1 + \beta) He$$

$$= (1 + \beta) P_w \cdot La \cdot E$$

substituting this value in eqn ③ we get

$$Hn - Hg - Hi - Hs = (1 + \beta) P_w \cdot La \cdot E$$

$$E = \frac{Hn - Hg - Hi - Hs}{P_w \cdot La \cdot (1 + \beta)}$$

E can be measured calculating

all the above H terms and calculating

β on eqn ②. For short periods

of less than a week, Hi and Hs

can be neglected as negligibly

small. The above eqn can be

written as

$$E = \frac{Hn - Hg}{P_w \cdot La \cdot (1 + \beta)} \quad (\text{for small periods})$$

Estimation of evaporation from a lake by the above method, gives satisfactory result, with error of the order of 5%.

These all other methods like, like wall-t-saucer method etc., to determine evaporation loss, but the most important widely used first method is pan measurement method.

Evaporation Reduction:

The different methods for reducing the evaporation from water surface are

(1) Reduction in surface area

(2) Mechanical covers

(3) Suppression of evaporation by surface films
↳ (Reducing)

(1) Reduction in surface area

The evaporation is reduced with the reduction (reducing) in surface area.

Evaporation is directly proportional to the exposed water spread area.

The different techniques for reducing the evaporation are as follows:

(i) By selecting conventional reservoirs sites.

(ii) By constructing reservoirs in minimum area which will maximum storage.

(iii) The water is stored below the ground. evaporation can be reduced.

(iv) Stream channels (streams) are straight and the evaporation is less.

(v) By storing water in one large reservoir rather than (reservoir) storing it in a number of small reservoirs, the evaporation can be reduced.

(2) Mechanical covers:

The mechanical cover is highly recommended in the case of small reservoirs.

(i) The mechanical cover while using reducing of evaporation.

(ii) Various types of mechanical covers are roofs, floating, drafts, wind break, floating light weight concrete block etc.

(iii) The loss of water due to transpiration

can be reduced by the removal of aquatic plants near the water surfaces.

(4) Suppression of evaporation by surface films
↳ (Reducing)

The molecules attach the water at one end and other end will not attached.

* These chemically act as evaporators & these chemically act as evaporators & suppress (reduce evaporation)

* This forming a layer on water surface

This monolayer of thickness less than 1 micrometre is formed on water surface

* which suppresses (reduces) the evaporation.

* This layer is invisible.

* the mono layer formed due to rain fall but it's formulated again after its breakdown.

* It is a flexible and moveable over the water surface.

The factors affecting the mono-layer

(a) Action of wind
(b) Oxidation

(c) Application with oil (HLL area, ocean)

The mono layer reduces the evaporation but don't reduce the amount of solar energy

(1) A reservoir with a surface area of 300 hectares had the following average meteorological values during a given week.

water temperature = 30°C

Relative humidity = 50%

wind velocity at 1m above ground = 12 km/h.

mean barometer reading = 750 mm of Hg.

Estimate the average daily evaporation from the lake reservoir and the volume of water evaporated from the lake during this week. Make use of Mayeot's formula as well as Rohwer's formula. Take saturation of vapour pressure at 30°C as 31.82 mm of Hg.

Sol:-

Mayeot's formula:-

\because barometer is an instrument measuring of atmospheric pressure

$$E = Km [es - ea] \left[1 + \frac{Vg}{16} \right]$$

1 mm of Hg = a unit

of pressure
= 0.001316
atmospheric
pressure units
force/unit
area

Given data

es = saturation of vapour pressure @ 30°C
= 31.82 mm of Hg.

$Km = 0.36$ for large deep water. $\rightarrow [\because$ From standard table,
7.23 SK.Garg]

$$ea = ?$$

$$\frac{ea}{es} = 50\% \quad (\text{relative humidity}) \quad [\because \frac{ea}{es} = RH]$$

$$\begin{aligned} ea &= 50\% es \\ &= 0.5 \times 31.82 \\ &= 15.91 \text{ mm of Hg} \end{aligned}$$

$$V_q = ?$$

$$V_1 = 12 \text{ km/h}$$

$$\frac{V_q}{V_1} = \left[\frac{9}{1} \right]^{1/7}$$

$$V_q = V_1 [9]^{1/7}$$

$$\therefore V_q = 12 \times [9]^{0.1113}$$

$$= 16.43 \text{ km/h}$$

\therefore wind velocity at
1m above ground
given in problem

\therefore The value of $V_q(0.8) V_{0.6}$ is
computed the following eq'n

$$V_Z = C \cdot Z^{1/7}$$

$$\therefore \frac{V_{Z_1}}{V_{Z_2}} = \left[\frac{Z_1}{Z_2} \right]^{1/7}$$

where
 V_Z = velocity at auf
height Z

substituting values we get

~~$E = 0.36 [3]$~~

$$E = 1 \text{ mm} [e_s - e_a] \left[1 + \frac{V_q}{16} \right]$$

$$= 0.36 [31.82 - 15.91] \left[1 + \frac{16.43}{16} \right]$$

$$= 11.61 \text{ mm/day}$$

Total evaporation valued in 7days (1 week) from 300 hectare
of surface area.

$$= \frac{11.61}{1000} \times 7 \times (300 \times 10^4) \text{ m}^3$$

$$= 243,810 \text{ m}^3$$

(b) Rohwer's formula:-

$$E = 0.771 (1.465 - 0.000732 P_a) (0.44 + 0.0733 V_{0.6}) (e_s - e_a)$$

where

$$e_s = 31.82 \text{ mm of Hg}$$

$$e_a = 15.91 \text{ mm of Hg}$$

$$P_a = 750 \text{ mm of Hg}$$

$$\frac{V_{0.6}}{V_1} = \left[\frac{0.6}{1} \right]^{1/7} \rightarrow \text{∴ wind velocity at } 1\text{m}$$

$$V_{0.6} = \cancel{V_1} [0.6]^{1/7}$$

$$= 12 [0.6]^{0.143}$$

$$= 11.15 \text{ km/h.}$$

$$E = 0.771 (1.1465 - 0.000732 \text{ Pa}) (0.4h + 0.0733 V_{0.6}) (e_s - e_a)$$

$$= 0.771 (1.1465 - 0.000732 \times 750) (0.4h + 0.0733 \times 11.15) \\ (31.82 - 15.91)$$

$$= 14.13 \text{ mm/day}$$

$$\text{Total evaporation} = \frac{14.13}{1000} \times 7 \times (300 \times 10^4 \text{ m}^2)$$

$$= \frac{14.13}{100} \underset{\text{day}}{\cancel{1}} \times \underset{7 \text{ day}}{\cancel{7}} (300 \times 10^4 \text{ m}^2)$$

$$= 2,96,730 \text{ m}^3$$

Conclusion:-

The ~~the~~ evaporation loss as computed by Rohwer's formula is about 20% higher than that calculated by Meynards formula. Hence, such formulas give only rough approximate values.

However, Rohwer's formula is generally preferred to all others.

UNIT-II Problems

(1) A reservoir with average surface spread of 4.8 km^2 in the first week of November had water surface temperature of 30°C and relative humidity of 40%. wind velocity measured at 3.0m above the ground is 18 km/h. The mean barometer reading is 760 mm of Hg. Calculate the average evaporation loss from the reservoir in mm/day and the total depth and volume of evaporation loss in the first week of November. Use both Meyer's equation as well as Rohwer's equation. Take saturation vapour pressure at 30°C as 31.81 mm of Hg.

Sol

(a) using Meyer's formula.

$$E = K_m (e_s - e_a) \left[1 + \frac{v_q}{16} \right]$$

Given that

$$e_s = 31.81 \text{ mm Hg}$$

$$\text{Relative humidity } R_H = 40\% \\ = 0.4$$

$$e_a = e_s \times R_H$$

$$= 31.81 \times 0.4$$

$$= 12.72 \text{ mm Hg}$$

$$v_q = v_2 \left[\frac{q}{3} \right]^{1/7}$$

$$= 18 \left[\frac{q}{3} \right]^{1/7}$$

$$= 18 \times 1.1699$$

$$= 21.06 \text{ km/h}$$

$K_m = 0.36$ for large ^{deep} water

$$E = 0.36 (0.4 - 12.72) \left[1 + \frac{21.06}{16} \right]$$

$$= 15.91 \text{ mm/day}$$

We know that

$$[\because \frac{e_a}{e_s} = R_H]$$

$$e_a = e_s R_H$$

∴ velocity is measured
at 3.0m

Total mean depth of evaporation in one week

$$= 7 \times 15.91$$

$$= 111.4 \text{ mm.}$$

Total volume of water evaporated = $(111.4 \times 4.8 \times 10^6) \times 10^{-3}$

$$= 0.5347 \times 10^6 \text{ m}^3$$

$$= 53.47 \text{ hectare-m.}$$

(b) using Rohwer's formula -

$$E = 0.771 (1.165 - 0.000732 P_a) (0.44t + 0.0783 V_{0.6}) (e_s - e_a)$$

Given that

$$e_s = 31.81 \text{ mm Hg.} \quad \text{and} \quad e_a =$$

$$e_a = 12.724 \text{ mm Hg.}$$

$$P_a = 760 \text{ mm Hg.}$$

$$V_{0.6} = \left[\frac{0.6}{2} \right]^{1/7} \times 18$$

$$= 15.16 \text{ km/h.}$$

$$E = 0.771 (1.165 - 0.000732 \times 760) (0.44t + 0.0783 \times 15.16)$$

$$\times (31.81 - 12.724)$$

$$= 20.74 \text{ mm/day}$$

∴ Total evaporation for one week = 20.74×7

$$= 145.2 \text{ mm.}$$

∴ Total volume of water evaporated = $(145.2 \times 4.8 \times 10^6) \times 10^{-3}$

$$= 0.697 \times 10^6 \text{ m}^3$$

$$= 69.7 \text{ hectare-m.}$$

Thus, we find that Rohwer's formula gives about more higher result than Mayers' formula.

EvaPo-tension Pidation

* EvaPo-tension Pidation is a combined term of evaporation and transpiration.

* TC is defined as the total loss of water through evaporation and transpiration from the plants.

* Transpiration is the process which water is lost through the living plants during the respiration process and back to the atmosphere.

* Evapotranspiration is more during summer season of greater amount of solar energy is received on the earth surface.

* The evapotranspiration rate falls down with increase in relative humidity which is most observed on rainy days.

Factors affecting evapotranspiration:-
The different factors which affect the evapotranspiration are as follows.

(1) climatic factors

(2) crop characteristics

(3) meteorological factors

(4) soil characteristics

(5) surface of leaves.

(6) Temperature

(7) Atmospheric conditions

(8) parameters of rainfall

(9) surface area.

(10) climatic factors

* The evapotranspiration is more during the month of summer. more amount of solar energy received in this period

* The solar radiation supplies energy for evapotranspiration process. Thus fall increases strength of the day increases the rate of evaporation

(a) crop characteristics:-

The evapotranspiration is high at seedling stage but decreases after the opening of seeds.
Hence higher the crop density, higher is the evaporation.

(b) meteorological factors:-

The full surface evaporation evapotranspiration is influenced by meteorological parameters. The parameters are the increase in temperature, sunlight and all attractive forces, increases the evapotranspiration but decrease the humidity.

(c) soil characteristics:-

Hydraulic conductivity and water holding capacity of soil effect evapotranspiration. The dry density of soil may decrease, evapotranspiration also decreases.

(d) surface of Leaves:-

Dark colored leaves (e) more greenish colored leaves will reflect the solar radiation, increase light colored surface of leaves reflect the more radiation. so evapotranspiration decreases.

(e) Temperature:-

The rate of evaporation more influenced by the temperature than any other factor.

(f) Atmospheric conditions:-

The evapotranspiration varies inversely with atmospheric humidity.

i.e. Atmospheric pressure increases evapotranspiration will be decreases.

At mean humidity will be more evapotranspiration will be decreases.

(8) Parameters of rain fall:-

The rate of evapotranspiration increases with increase in frequency and amount of rain fall.

(9) Surface area:-

In irrigated area which all surrounded by large area (very dry) (or) semi-arid (give dry) condition. The rate of evapotranspiration is more.

Measurement of evapo-transpiration.

Measurement of evapotranspiration can be measured in two ways

(1) Direct measurement

(2) Measurement by use of evaporation

(1) Direct measurement:-

The direct measurement of evapotranspiration can be measured by five principal methods

(a) Tank lysimeter methods

(b) Field experimental plots

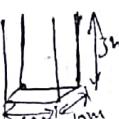
(c) Soil moisture studies

(d) Irrigation methods

(e) Inflow and out-flow studies for large area

(a) tank lysimeter methods:-

* Tank container with ground level having an area of $10m^2$ and 3m deep



* Large size of tank is useful to the greater soil development.

* A tank is filled with soil of the field and crop is grown in it.

* To measure the required quantity of water and to maintain the constant moisture condition of the crop with the tank proper growth

(b) field experimental plots:-

* In this method irrigation water is applied to the selected fields in such a way that there is neither run off nor deep percolation.

* It is seen from observations the more amount of water will be used. The yield of the crop will increase rapidly but in certain point it will then decrease with further increase in water.

* That is break down in the curve and that is consumptive use of water in the field.

(c) soil moisture studies:-

The soil moisture is measured before and after the each irrigation.

* The quantity of water (observed) per day from soil is computed for each period.

* A curve is drawn by plotting the rate of use of water against time and this curve

(d) Integration method:-

The integration method is summation of the products

~~per unit~~

(i) ~~per~~ unit consumptive use for each crop times its area.

(ii) unit consumptive use of native vegetation times its area.

(iii) water surface evaporation times ~~area~~ surface area

(iv) evaporation from bare land times its area.

→ You take a some area will be irrigated area on that area crop is grown the period of 6 months that crop after that land will not use so soil will effect evaporation so on that time crop the another seed ~~soil~~ means soil will be no loss of moisture

(a) Inflow-outflow studies for large areas

In this method, annual consumptive use is calculated for large areas.

$$U = (I + P) + (G_S - G_E) - R$$

where

U = Valley Consumptive use.

I = Total inflow during 12 months (year)

P = Yearly Precipitation

G_S = Ground storage at the beginning of the year

G_E = Ground storage at the end of the year

R = Yearly out-flow.

(2) Measurement by use of equations:-

(a) Penman's equation.

(b) Blaney-Criddle formula.

(c) Hargreaves Class A Pan evaporation method.

(a) The Penman's equation:-

The Penman's equation is used to calculate the potential evapotranspiration by combining the energy balance and mass balance methods.

* According to Penman, The Potential evapotranspiration PET is given by

$$PET = \frac{AH + EA_R}{A + T}$$

where $A = \text{slope}$

$\gamma = \text{psychrometric constant}$
(0.00656 of m/mg/°C)

$H = \text{Net radiation}$

$$H = Ha \left(1 - \left[a + \frac{bn}{N} \right] \right) - \sigma (T)^4 (0.56 - 0.092) \sqrt{0.10 + 0.9 \frac{n}{N}}$$

$Ea = \text{mass transpiration}$

$$= 0.35 \left[1 + \frac{U_2}{T_{10}} \right] (e_s - e_a)$$

(b) Blaney-Criddle formula:-

Blaney-Criddle has given the accurate empirical formula which daily Potential evapotranspiration (PET) can calculate.

$$PET = 2.5 K F$$

where $K = \text{Empirical coefficient}$

$$F = \frac{PT^4}{100}$$

$P = \text{mean daily percentage}$

$T = \text{mean daily temperature (t)}$

$$T = \frac{T_{\text{maximum}} + T_{\text{minimum}}}{2}$$

(b)

(c) Hargreaves (Class A Pan evaporation method)

This method is very much used in India, after from its use in USA, Philippines and U.S.A according to Tuin method the consumptive use (U) (or) Evapo-transpiration (E_T) is given by

$$E_T (or) U = K E_P$$

where $K = \text{consumptive use coefficient on different crops}$

$E_P = \text{Class A pan evaporation}$

$$E_P = 0.456 R_A \cdot C_T \cdot C_H \cdot C_S \cdot C_E (in)$$

where $R_A = \text{extra-terrestrial radiation (m)}$

$C_T = \text{coefficient of temperature.}$

$C_H = \text{coefficient of wind velocity}$

$C_S = \text{coefficient of possible sunshine}$

$C_E = \text{relative humidity}$

$C_E = \text{elevation}$

$$C_T = 0.393 + 0.02796 T_C + 0.0001189 T_C^2$$

$$C_W = 0.708 + 0.0034 W - 0.0000038 W^2$$

$$C_S = 0.542 + 0.10085 - 0.78 \times 10^{-4} T_C + 0.62 \times 10^{-6} T_C^2$$

$$h = 1.250 - 0.0087H + 0.75 \times 10^4 H^2 - 0.83 \times 10^{-8} H^4$$

$$I_e = 0.97 + 0.00984E$$

(E being elevation in 100 m res)

Infiltration :-

Infiltration may be defined as the downward movement of water from soil surface into the soil mass through the pores of top soil.

In simple terms infiltration is the entry (or) passage of water onto the soil through the soil surface. Once water enters onto the soil, the process of transmission of water known as percolation.

Infiltration and percolation are directly inter-related when percolation stops.

Infiltration stops (white head) this topic affecting infiltration factors

(1) Condition of entry surface : (vegetation covers, versus bare land)

(2) Permeability

(3) Antecedent moisture condition on soil.

(4) Temperature.

(5) Intensity and duration of rainfall.

(6) Movement of man and animals

(7) Change due to human activities

(8) Quality of water

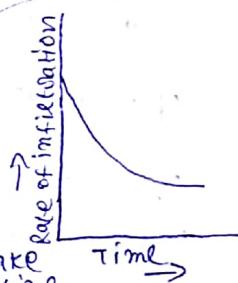
(9) Presence of ground water table.

(10) Size and characteristics of soil particles.

(11) Catchment parameters.

Infiltration capacity curve :

Any soil to absorb the water from rain fall once the soil absorbs the minimum amount of water after that soil absorb maximum water slowly take more time



(1) Condition of entry surface:-

If the area covered with grass, vegetation and bushy plants, infiltration capacity will be more.

The rain drops falling on the surface will cause in-washing of fine particles of the soil and will clog (block) the surface thus resulting in the reduction (reduction) of infiltration.

The vegetation covering removes soil moisture through transpiration. The rate will be more, and the infiltration rate will be more.

(2) Permeability :-

Permeability means permit the water inside the soil mass.

Clayey cotton soils hardly permit and loamy soils some more permeable, sandy soils are more permeable.

The permeability or infiltration of water on soil depends upon the several factors such as type of particle size, presence of organic matter (dead plants) present in the soil.

(3) Antecedent moisture condition on soil.

The moisture condition present on the soil (it means if the soil is wet soil) infiltration rate will be less. The soil is dry soil infiltration rate will be high.

Because the heat will be more on soil. It means temperature will be more the soil will expand temperature will be less the soil will be contracted.

(1) Temperature:-

The viscosity (viscousity) of water changes with temperature.

In summer therefore infiltration will be higher due to the less viscous water and also particle-to-particle size of earth particle is elongated so infiltration will be more. In winter therefore infiltration will be lower due to the viscous water and also particle-to-particle size of earth particle

is compression stage so there is no gaps of earth particles so infiltration rate will be lower.

(2) Intensity and duration of rainfall

When the intensity of rainfall will be more will be falling in less area. This water will be compacted to the earth surface and will wash off fine particles, so gaps of the soil particles will be filled with soil resulting infiltration decrease in the rate of infiltration.

so less intensity of rain fall will be given more infiltration rate, higher intensity of rain fall will be given less infiltration

⑦

(3) Movement of man and animals

When the heavy movement of man or animals on soil surface, the soil will be compacted resulting gaps of soil particle to particle will be decrease so reduce the infiltration rate

(7) change due to human activity:-

cultivation of barren land, by growing crops and grass cover, resulting increased rate of infiltration. On the other hand, construction of roads, houses, factories, play grounds, over-grazing (eating) of pastures (animal result in reduction in infiltration capacity).

(8) quality of water:-

silt, and other impurities are present in incoming water result reduction of infiltration rate due to lack of closing of soil particle gaps

The salts are present in water viscosity (viscous) will be less and this water stick with soil to form completely that reduce the infiltration rate

(9) presence of ground water table:-

→ The ground water table will ~~be~~ draw at the top soil surface in this case infiltration rate will be less
→ The ground water table will be ~~be~~ depth of ground in this case infiltration rate will be more

(10) size and characteristics of soil particle:-

The size of soil particles is bigger size result, particle to particle will be more result, more infiltration
The size will be less of particle to particle result gaps will be less result, less infiltration infiltration rate is directly proportional to the grain size/diameter of granular soils.

In other words the soil has swelling minerals like illite and montmorillonite.

The infiltration rate will be reduced.

(M) Catchment parameters:-

Catchment area will be drainage. The water runoff will be due to water will be moving one place to another place and stored in the place so the water will be runoff, so infiltration rate will be less.

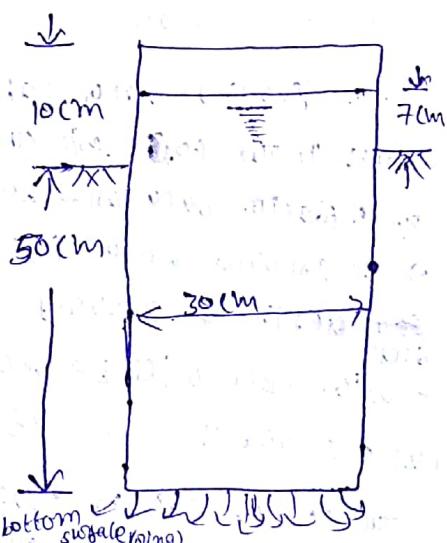
Catchment area will be increase the running of water will be decrease so infiltration rate will be increased.

Filled measurement of infiltration rate:

Infiltration is to be measured in the field in two ways.

- single tube infiltrometer.
- double tube (or double ring) infiltrometer

(a) single tube infiltrometer:-



→ It consists of hollow meter cylinder of 30 cm diameter and 60 cm length and both ends open. It is a measuring device which is used for measuring infiltration rate.

→ The cylinder is driven into the ground at a level of such that 10 cm of it projects above the ground.

→ The cylinder is filled with water upto a head of 7 cm (i.e. means 57 cm filled with water)

→ The infiltration meter is maintained above the ground level.

→ Due to the infiltration of water slowly the water level on the cylinder will be decreasing.

→ The decreasing of water will be added to the cylinder will be maintained. The constant level the added water will be measuring.

level by jar or burette.

→ The observations are taken about 3 to 6 hours depending upon the type of soil.

→ The major drawback of single tube infiltrometer is that the infiltrated water percolates laterally at the bottom of the ring of shown 30 cm.

→ Hence the bottom of the soil surface is saturated having some moisture content due to absorbing the water. This moisture content will be absorbed by the both sides of the soil in the tube, so the tube below soil is dried soil.

so infiltration rate will be more so infiltration rate will be more. Hence it does not truly represents the soil how much area is saturated in the bottom surface.

→ Hence it does not truly represents the soil how much area is saturated in the bottom surface.

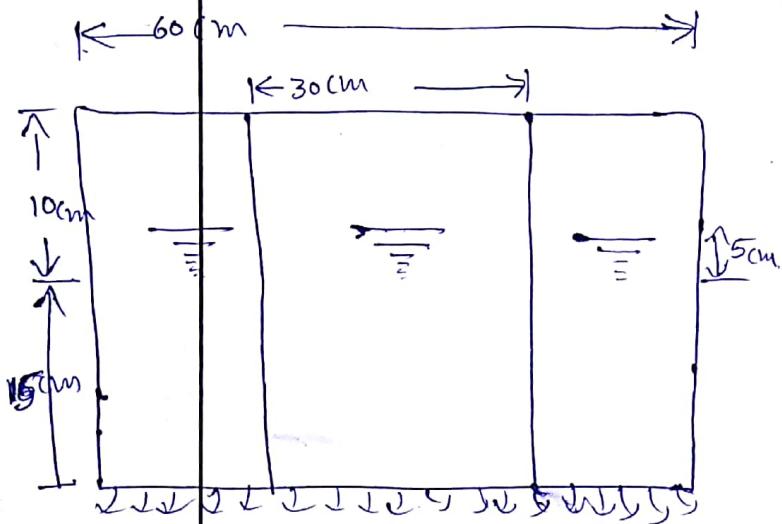
→ Infiltration rate is determined.

→ Infiltration rate is determined by the following formula.

(b) Double tube infiltrometer:-

The previous defect of single tube infiltrometer while rectifying of this problem two types double tube infiltrometer will be introduced

- * In double tube infiltrometer consists of two concentric hollow rings (or cylinders) driven in to the soil uniformly without any tilt and disturbing the soil to the depth of 15cm
- * The diameter of the ring may vary from 25 to 60 cm



- * The purpose of introducing into pressure the lateral percolation of water is applied to the water in the soil
- * The soil will be saturated and will not absorb the moisture content in middle tube of soil moisture
- * The observations are continued till constant infiltration rate is observed (measured in ~~in~~ 1 hour 5 cm of water infiltrated and next 1 hour also 5 cm of water is infiltrated that is constant infiltration)

- * The water is applied in both inner and outer rings to maintain a constant depth of about 5 cm
- * The water is slowly will be infiltrated absorb the ground, the 1 cm of water will be decreasing on to the tube after that the 1 cm of water will be added. The added water will be measured by using of measuring jar (or bucket)
- * The water depth in the inner and outer rings should be kept the same during the observation period.

UNIT-II : Problems

162

163

P-332 (8K-Barr)

- (3) A reservoir with a surface area of 300 ha (long) had the following avg. meteorological values during a given week.

$$\text{water temperature} = 30^\circ\text{C}$$

$$\text{relative humidity} = 50\%$$

$$\text{wind velocity at } 1\text{ m above ground} = 12 \text{ km/h}$$

$$\text{mean barometer reading} = 750 \text{ mm of Hg}$$

Estimate the avg daily evaporation from the lake reservoir and the volume of water evaporated from the lake during this week. Make use of Meyer's formula as well of Rohwer's formula to compare the results.

Sol: Meyer's formula.

$$(i) E = km [es - ea] \left[1 + \frac{V_q}{T} \right]$$

{: barometer is a instrument measuring of atmospheric pressure}

Given: $e_s = \text{saturation vapour pressure at } 30^\circ\text{C}$

$$= 31.82 \text{ mm of Hg} \rightarrow \text{(from std table 7.23)}$$

$$km = 0.36 \text{ for large deep water} \rightarrow \text{sk. Georg}$$

$$ea = 8$$

$$\frac{ea}{es} = 50\% \text{ (relative humidity)}$$

~~$$ea = 0.5 es$$~~

$$= 0.5 \times 31.82$$

$$= 15.91 \text{ mm of Hg}$$

$$V_q = ?$$

$$V_1 = 12 \text{ km/h}$$

$$\frac{V_q}{V_1} = \left[\frac{q}{1} \right]^{1/7}$$

$$V_q = V_1 \cdot \left[\frac{q}{1} \right]^{1/7}$$

{: wind velocity
= 12}

$$v \cdot v_a = [9]^{0.163} \times 12 \text{ km/h}$$

$$= 16.43 \text{ km/h}$$

Substituting values, we get

$$E = 0.36 [31.82 - 15.91] \left[1 + \frac{16.43}{16} \right]$$

$$= 11.61 \text{ mm/day}$$

Total evaporation value in 7 days (1 week) from 300 hectares

of surface area

$$= \frac{11.61}{1000} \times 7 \times (300 \times 10^4) \text{ m}^3$$

$$= 2,438,10 \text{ m}^3$$

(b) Rohwer's formula

$$E = 0.771 (1.465 - 0.000732 P_a) (0.44 + 0.0733 V_{0.6})$$

where

$$e_s = 31.82 \text{ mm of Hg}$$

$$e_a = 15.91 \text{ mm of Hg}$$

$$P_a = 750 \text{ mm of Hg}$$

$$V_{0.6} = \left[\frac{0.6}{T} \right]^{1/2} \quad \therefore \text{wind velocity} = 1$$

$$V_{0.6} = [0.6]^{0.163} \times 12 \text{ km/h}$$

$$= 11.15 \text{ km/h}$$

$$E = 0.771 (1.465 - 0.000732 \times 750) (0.44 + 0.0733 V_{0.6})$$

$$(31.82 - 15.91)$$

$$= 11.13 \text{ mm/day}$$

$$\begin{aligned}
 \text{Total evaporation} &= \frac{16.13}{1000} \times 7 \times (300 \times 10^4 \text{ m}^2) \\
 &= \frac{16.13}{100} \frac{\text{m}}{\text{days}} \times 7 \text{ days} \times (300 \times 10^4 \text{ m}^2) \\
 &= 2,96,730 \text{ m}^3
 \end{aligned}$$

Conclusion-

The evaporation day as computed by Roth's formula is about 80% higher than that calculated by Moyer's formula.

Hence, such formulas give only rough approximate values.

However, Roth's formula is generally preferred to all others.

- Ques No 38 (SKG 88)
- 2) A storm with a 15.0 cm precipitation produced a direct runoff of 8.7 cm. The time distribution of the storm

- ① runoff of 8.7 cm. The time distribution of the storm

is as follows

Time (in hours) vs. Rainfall (in cm) per hour

Time from start in hr	1	2	3	4	5	6	7	8
Incremental rainfall in each hr in cm	0.6	1.35	2.25	3.15	2.7	2.4	1.5	0.75

Estimate the phi index of the storm.

